

Regional contrasts in surface pCO₂ across Northern, Central and Southern Bay of Bengal

Louwin Anand D^{1*}, Ayyapan Saravankumar¹

CAS in Marine Biology, Annamalai University, Faculty of Marine Sciences,
Cuddalore, Tamil Nadu, India

Correspondence: louwinanand@gmail.com

Abstract

The Bay of Bengal (BoB) regulated by monsoonal seasons, driven by monsoonal winds, hydrography and riverine inputs responsible for the carbon equilibrium in ocean. In the present study, spatio-temporal variation of pCO₂ dynamics across Northern, Central and Southern BoB region assessed during the years 2022-2024. From the pCO₂ generated image, Northern BoB exhibited with higher pCO₂ values (upto ~450 μatm) compared to Central and Southern regions. Post-monsoon and pre-monsoon resulted with pCO₂ enriched season attributed to the huge load of freshwater-discharge via Ganges-Brahmaputra system, enhanced solubility under reduced salinity and monsoonal upwelling which aids in transport pCO₂ enriched subsurface water to the surface. In contrast, summer season resulted with low pCO₂ levels (~221 to 285 μatm), mainly responsible due to elevated Sea Surface Temperature (SST), thus facilitates strong thermal stratification and inhibits the vertical mixing. The Central and Southern BoB regions exhibited more low and stable values mainly driven by East Indian Coastal Current (EICC) and South-west Monsoon current (SMC). The observed findings highlight the collaborative role of physical, chemical and biological process shaping the pCO₂ variability in regional carbon dynamics. The long-term monitoring is adequate to understand the global carbon cycle under climate change conditions.

Keywords: pCO₂ variability, CO₂ solubility, riverine flux, EICC, upwelling

1.Introduction

The ocean acts as ocean sink basin for about 25% of anthropogenic CO₂ annually (Sabine et al., 2004, Takahashi et al., 2009; Le Qu' er' e et al., 2018). This sequestering process results with the CO₂ levels present in the atmosphere (Sabine et al., 2004). Since the ocean sequester a atmospheric CO₂, the CO₂ exchange varies with space and time. The air-water CO₂ flux (fCO₂) is primarily governed by the CO₂ transfer velocity (k) and surface partial pressure of CO₂ (pCO₂), together determines the efficiency of source and sink of capacity (Liu et al., 2022; Shen et al., 2023). The exchange among the CO₂ solubility pump, carbonate and biological pump can further modulate the conversion process of Dissolved Inorganic Carbon (DIC), thereby drive the sources and sink in aquatic ecosystem (Akhand et al., 2022).

The increased rate of acidification near coastal region reported to the various anthropogenic activities that modifies the carbonate chemistry in aquatic ecosystem (Doney, 2009; Cai et al., 2011,2020; Hall et al., 2020). Several factors which are responsible for the ocean acidification are atmospheric pollutants deposition via sulphate and nitrate aerosols attributed to the fossil fuel burning specifically in urbanized regions (Mackenzie et al., 1995; Schlesinger, 1997; Doney et al., 2007). Several recent studies reported that continental shelf considered as the sink basin (Laruelle et al., 2018). The south China sea, considered as the weak source (Dai et al., 2013, Zhai et al., 2013) and eventually turned out to be the atmospheric CO₂ sink (Li et al., 2020). Although understanding the carbon dynamics and assessing surface pco₂ trends relatively vital due to the lack of in-situ pCO₂ measurements, which possess the challenge in pco₂ estimation (Borges, 2005, Anderson, 2005, Anderson, 2010).

The present study explores the spatial variability of fCO₂ in the BoB region. The BoB region found unique due to its seasonal reversing and semi enclosed area with an opening in the south (Shetye., 1996). The regions receive higher freshwater influx from major riverine inputs from Ganges-Brahmaputra (Unesco, 1969) and heavy precipitation. The BoB region experiences two distinct regions including Southwest Monsoon (June- September) and Northeast Monsoon (December - February) driven by north-westerly winds. Along with the monsoon and major river system, delivers the major freshwater input in BoB, thus influence the physical characteristics hence altering the carbon dynamics (Joshi et al., 2021). The study reported by Muraleedharan and Prasannakumar, (1996), suggests that the BoB coast acts as atmospheric CO₂ sink due the increased biological productivity facilitated by the nutrients derived from river-runoff, thus diminishes the surface pCO₂ values. Notably, the rate of pCO₂ increase along with southwestern BoB coast align with global average of approximately 1.5uatm.yr⁻¹, meanwhile, northern coast demonstrated 3-5 times higher at 6.7uatm.yr⁻¹ (Sarma et al., 2015). The accelerated increase in northern BoB associated with increased deposition of sulphate and nitrogen aerosols during winter and spring months. The seasonal reversing coast currents such as East Indian Coastal Current (EICC) and West Boundary Current (WBC) responsible for the variations of hydrographic characteristics (Sarma et al., 2019).

Several studies examined the pCO₂ variability in BoB region across space and time. The present study aims to investigate the pCO₂ distribution with spatial and temporal variation in BoB regions subdivided into Northern, Central and Southern BoB regions with

relation to SST and chl-a by integrating satellite images. This provides new insights into the carbonate dynamics of this climatically significant region and its role in regional and global carbon budgets.

2. Material and Methodology

The satellite derived pCO₂ maps were generated using the empirical algorithm proposed by Shanthi et al, (2016). Monthly SST and chl-a products were obtained from MODIS-Aqua satellite datasets. All graphical representation and visualizations were demonstrated using R studio. For regional assessment, the BoB region was sub divided into Northern, Central and Southern sectors to evaluate the pCO₂ variability in relation to distinct hydrographic and biogeochemical characteristics of subregions.

pCO₂ algorithm

$$pCO_2 = 1025.682 - 7.7794 * SST + 6.0874 * chl - 0.5777 * SST^2 + 19.9015 * chl^2$$

Where, pCO₂= Partial pressure of Carbon-di-oxide, SST = Sea Surface temperature, chl = Chlorophyll concentration

The BoB region is sub categorized in to Northern, Central and Southern BoB based on the study carried out by Feng et al, (2022) and Cui et al, (2024), shown in Fig.1.

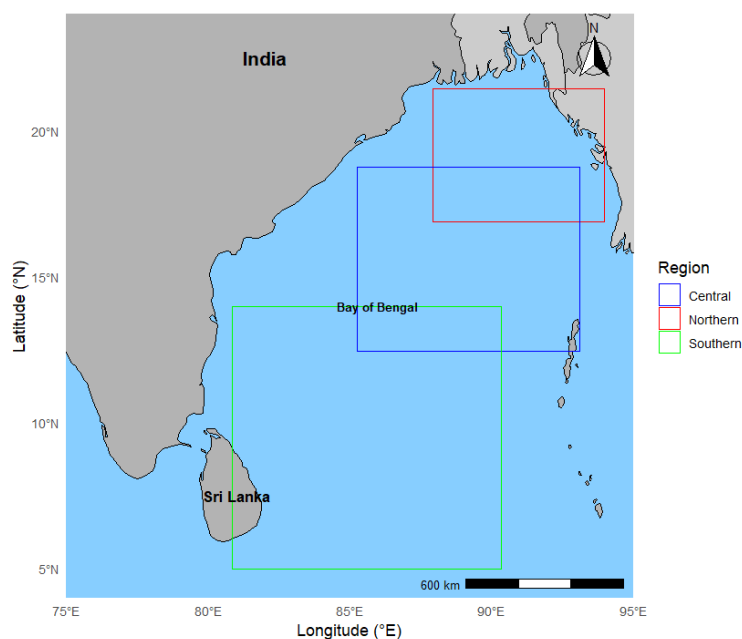


Fig.1 Map showing study area

3.Results

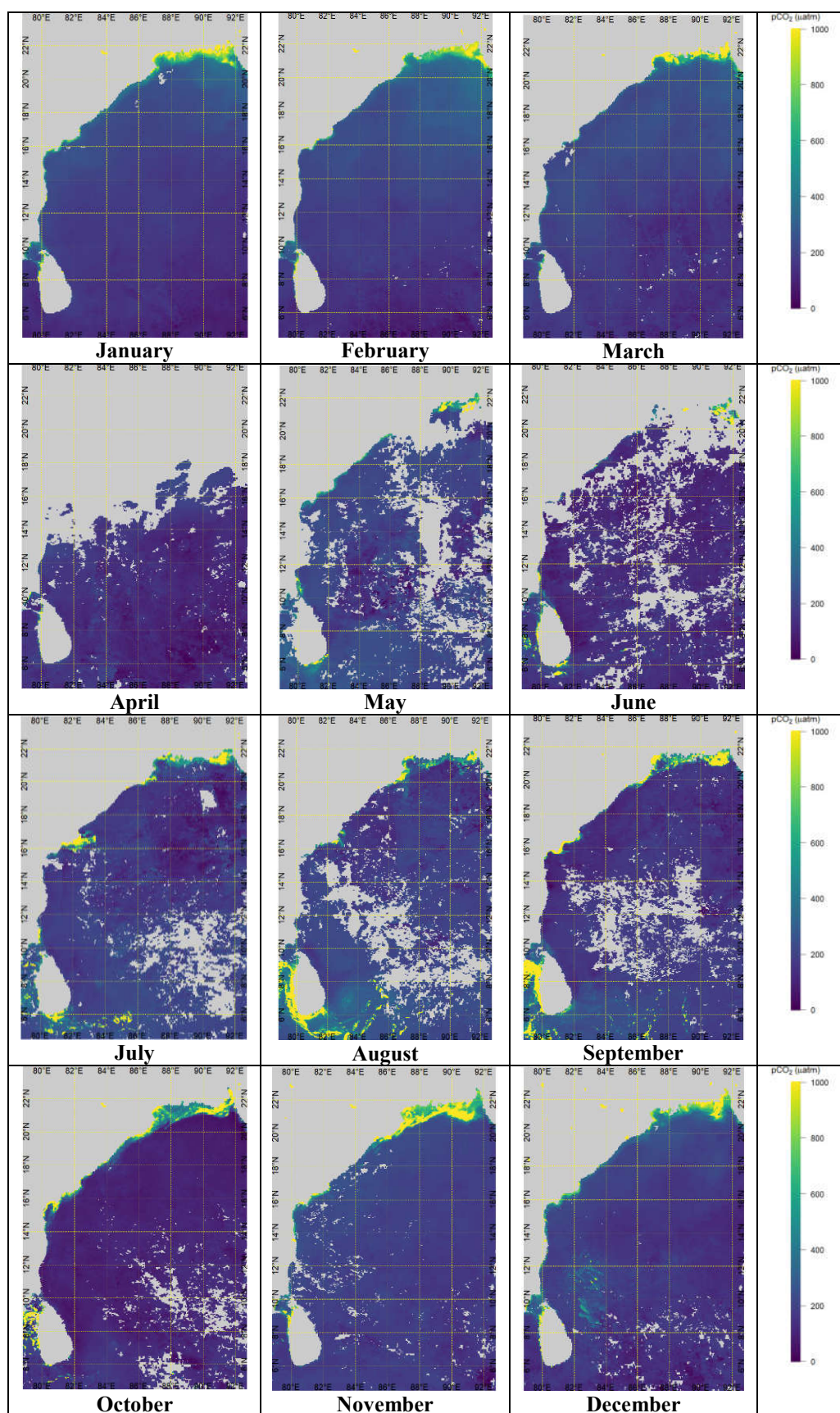


Fig.2 Monthly $p\text{CO}_2$ variation during the year 2021

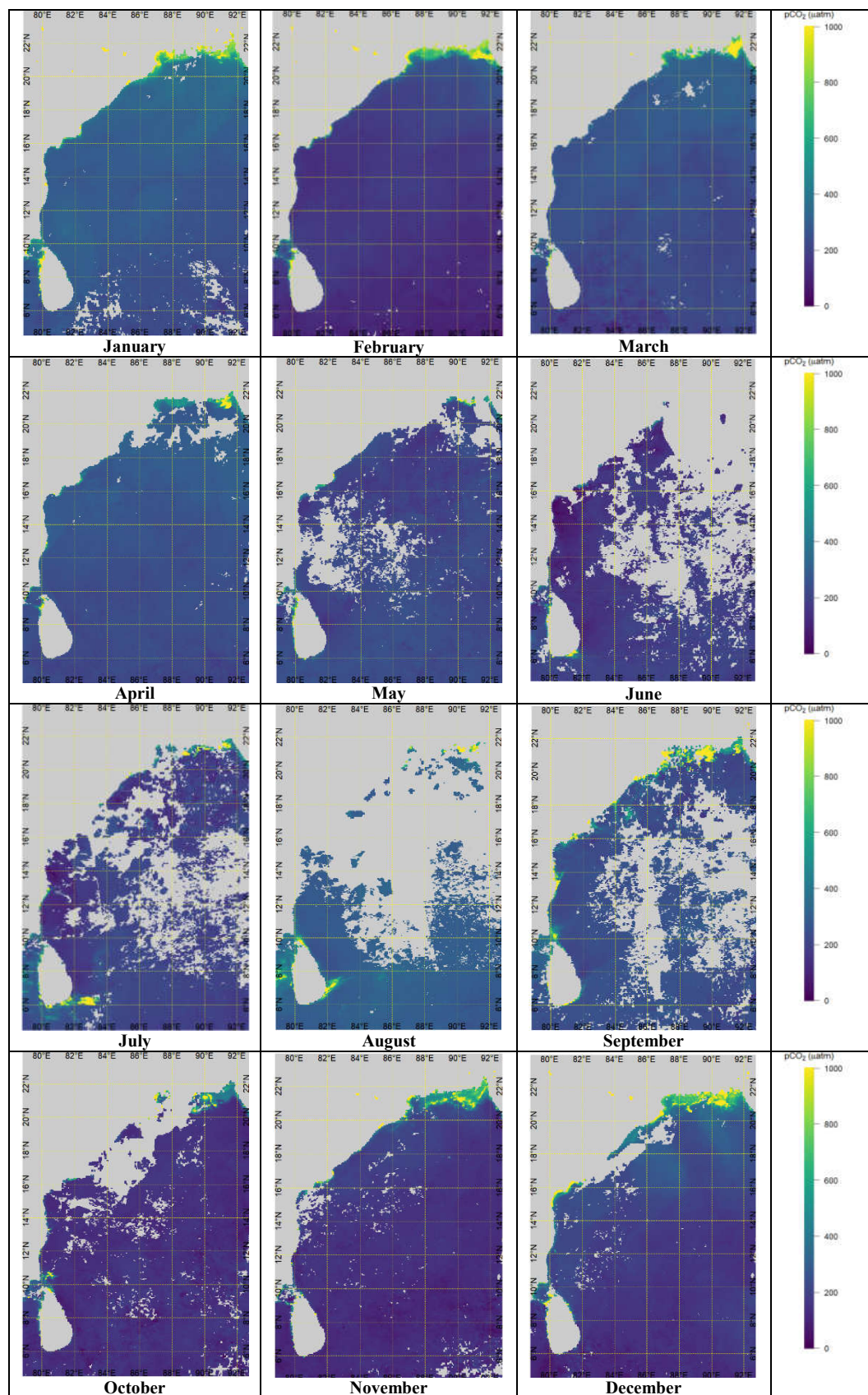


Fig.3 Monthly pCO_2 variation during the year 2022

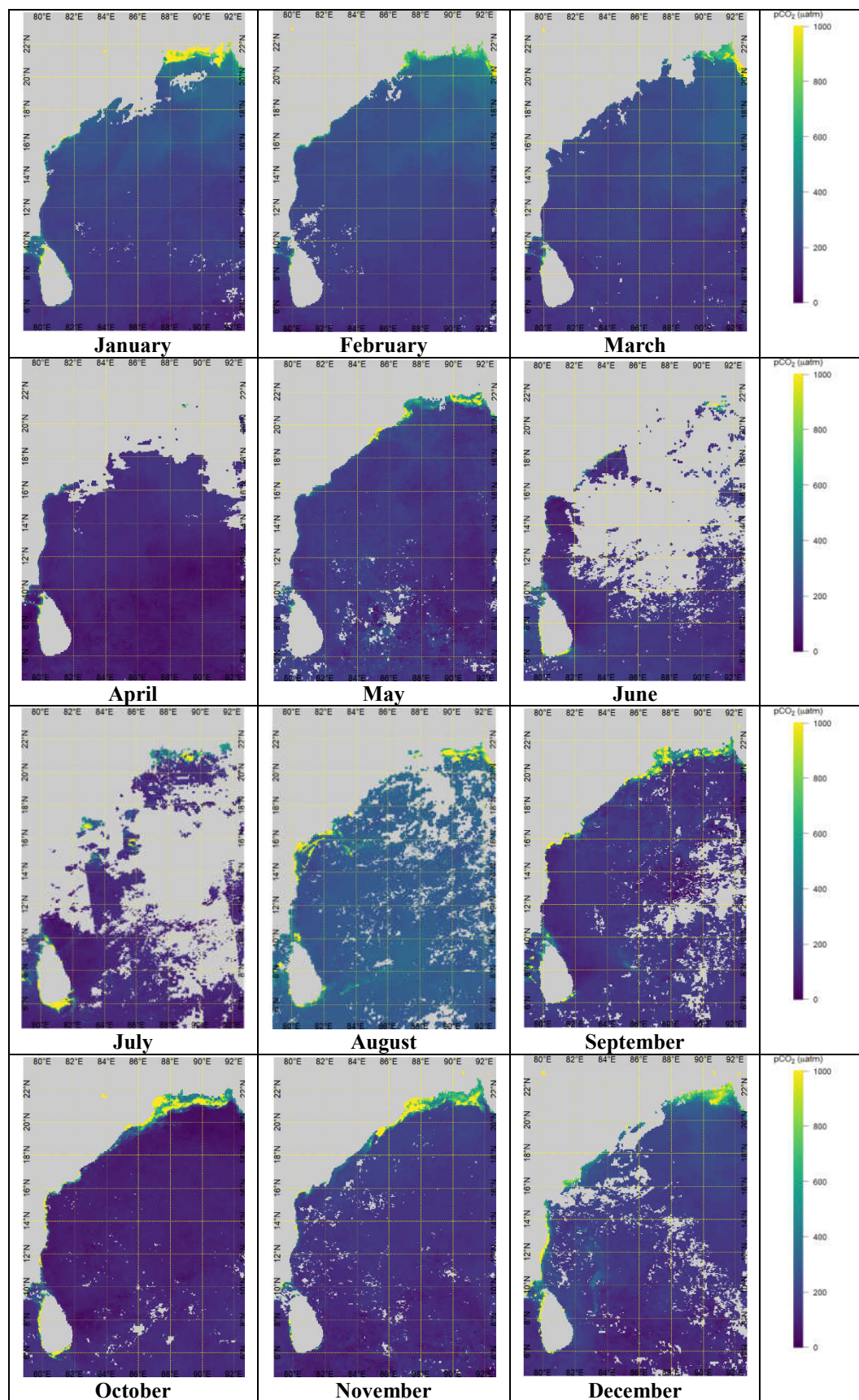


Fig.3 Monthly $p\text{CO}_2$ variation during the year 2023

3.1. Seasonal Variability of pCO₂ in the Bay of Bengal (2022–2024)

The seasonal distribution of pCO₂ across the Northern, Central, and Southern Bay of Bengal revealed distinct spatiotemporal patterns over the three-year period from 2022 to 2024. Across all three years, peak pCO₂ concentrations were consistently observed during the Pre-Monsoon and Post-Monsoon (Po-Monsoon) seasons, while relatively minimum values were recorded during the Summer and Monsoon periods.

3.1.1. Northern BoB

In 2022, Pre monsoon and Post monsoon registered higher pCO₂ levels 457.47 μatm (January) and 410.99 μatm (September), meanwhile minimum concentration observed in Monsoon (283.54 μatm, October) and Summer (294.34 μatm, May). In 2023, The minimum pCO₂ concentration noted were 274.68 μatm (June, Summer), 275.62 μatm (May, Summer) and in summer (274.68 μatm, June). In 2024, post-monsoon (488.78 μatm) and pre-monsoon (411.23 μatm) exhibited higher pCO₂ levels, meanwhile minimum pCO₂ levels noted are 248.75 μatm (May, Summer) and 307.31 μatm (October, Monsoon).

3.1.2. Central BoB

During 2022, low pCO₂ conditions observed during summer and monsoon are 244.62 μatm (April, summer) and 260.21 μatm (October, Monsoon). The peak pCO₂ levels observed in 392.26 μatm (February, Post-monsoon) and 351.89 μatm (September, Pre-monsoon). In 2023, 258.16 μatm and 293.12 μatm are the minimum pCO₂ levels observed, meanwhile, maximum pCO₂ levels registered during 394.4 μatm (January, post-monsoon) and 364.94 μatm (September, Pre-monsoon). During 2024, 388.63 μatm (January, Post-monsoon) and 346.08 μatm (December, Monsoon) noted as maximum pCO₂ levels and the other hand 215.88 μatm (May, Summer) and 272.93 μatm (October, Monsoon).

3.1.3. Southern BoB

In 2022, the low pCO₂ conditions observed are 290.33 μatm (March, Po-monsoon) and 307.68 μatm (July, pre-monsoon) and maximum pCO₂ levels are 342.55 μatm (January, po-monsoon) and 359.93 μatm (September, pre-monsoon). In 2023, 240.06 μatm (April, summer), 285.65 μatm (August, pre-monsoon) and 296.87 μatm (November, monsoon) are the maximum pCO₂ levels noted meanwhile, maximum pCO₂ levels registered are 372.89 μatm (January, po-monsoon) and 322.88 μatm (September, pre-monsoon). The minimum pCO₂ levels noted during 2024, 217.51 μatm (May, summer) 275.62 μatm (November, Monsoon) and the maximum pCO₂ levels are, 339.51 μatm (January, Po-monsoon) and 317.78 μatm (August, pre-monsoon).

Across all regions Po-Monsoon and Pre-Monsoon periods were marked by elevated pCO₂ levels, likely driven by enhanced remineralization, decreased biological uptake or physical mixing. Summer and Monsoon seasons noted a decline trend in pCO₂, possibly due to higher primary productivity and dilution from freshwater input via precipitation during monsoonal seasons. The Northern BoB exhibited the highest seasonal amplitudes, while the Southern BoB remained relatively stable.

These seasonal variations underscore the dynamic nature of carbon dynamics in the BoB, suggests that biological activity and monsoonal forcing play dominant roles in shaping surface pCO₂ distributions.

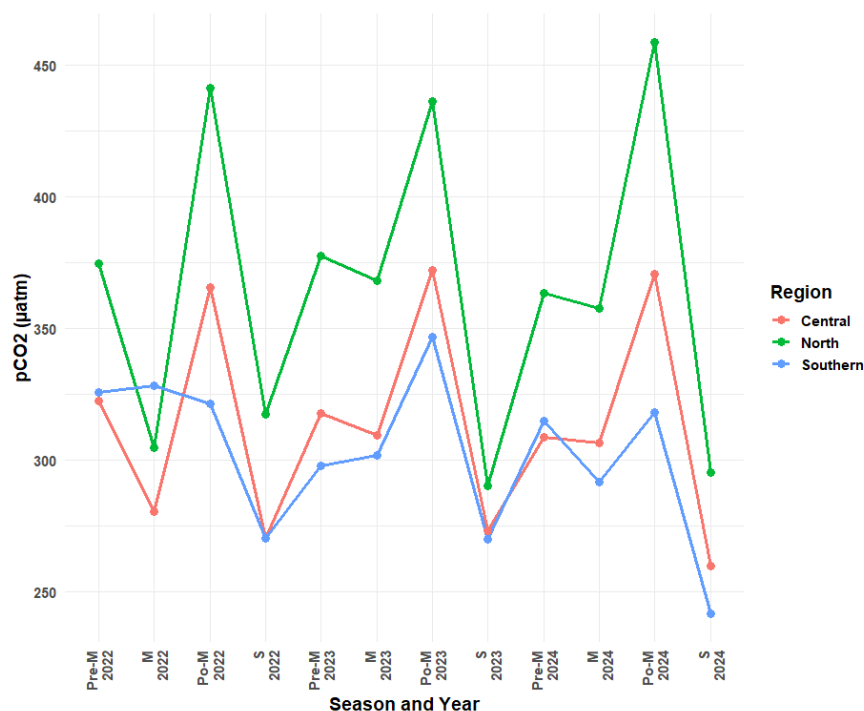


Fig.5 seasonal pCO₂ variation during the year 2021-2023

3.2. Seasonal and Spatial Variation of pCO₂ based on the heatmap correlation (2022–2024)

The spatial and seasonal patterns of pCO₂ across BoB regions exhibited distinct seasonal and inter-annual variations during the years 2022-2024.

In 2022, northern BoB regions exhibited highest seasonal pCO₂ values, peaking during the post-monsoon season (441.2 µatm), while lower pCO₂ values were observed during summer (317.6 µatm). The central BoB exhibits peak pCO₂ values in post-monsoon season (365.4 µatm) with diminished pCO₂ levels noted during summer (270.5 µatm). In contrast, southern BoB exhibits relatively stable values across seasons, although higher pCO₂ conditions during pre-monsoon (325.9 µatm) and lowest during summer (270.3 µatm), represented in Fig.6(a).

In 2023, a similar trend was observed, with Northern BOB exhibited a maximum pCO₂ levels during post-monsoon (436.2 µatm) and reduced levels during summer (290.3 µatm). The central BoB exhibited a similar pattern where peak pCO₂ levels noted during post-monsoon (372.1 µatm). while, lowest during summer (273.5 µatm). Southern BoB too exhibited a constant pattern registering minimum pCO₂ levels during summer (269.9 µatm) and maximum during post-monsoon season (346.8 µatm), demonstrated in Fig.6(b).

In 2024, post-monsoon season found consistent with peak pCO₂ levels in Northern (458.5 µatm), central (370.6 µatm) and Southern BoB (318.1 µatm). Meanwhile, summer

values with constantly low with Northern (295.6 μatm), Central (259.9 μatm) and Southern BoB (241.7 μatm), shown in Fig.6(c).

Throughout the study, post-monsoon season found to be dominant across the BoB and summer with poor pCO_2 levels. Among the regions, Northern BoB regions found consistently high followed by Central and Southern BoB regions. The dominant pCO_2 levels in Northern BoB region attributed to the influence of major riverine inputs and stratification. On the other hand, reduced pCO_2 levels in southern region indicates reduced mixing leads to the diminished carbon input.

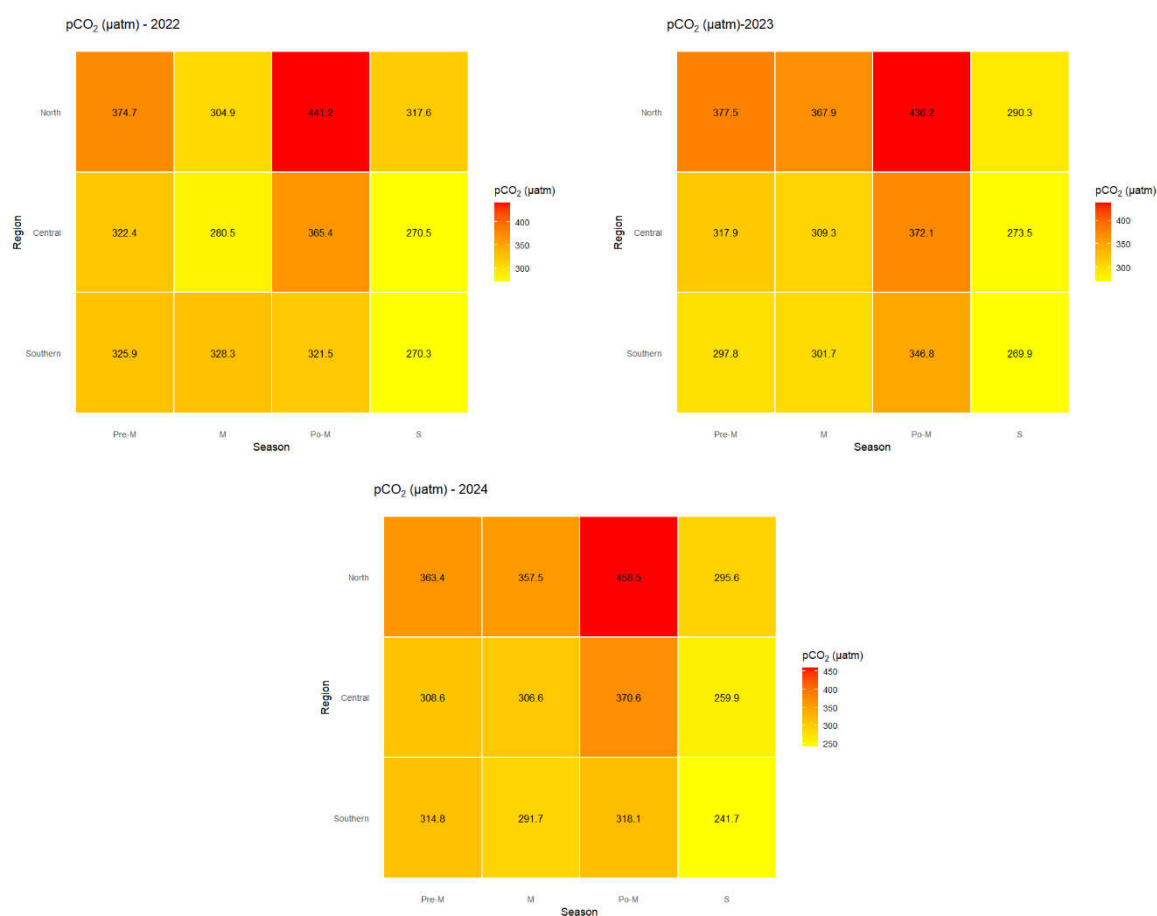


Fig.6 heatmap of pCO_2 variability (a) 2021, (b) 2022 and (c)2024

4. Discussion

The uptake of atmospheric CO_2 by the surface ocean is primarily regulated by the air–sea CO_2 flux. This flux is controlled by the partial pressure gradient of CO_2 between the ocean and atmosphere, together with environmental factors such as sea surface temperature (SST), sea surface salinity (SSS), and gas transfer velocity, the latter being strongly

influenced by wind speed (Wanninkhof, 1992, 2014; Nightingale et al., 2000; Ho et al., 2006; Wanninkhof et al., 2009; Blomquist et al., 2017). In the present study, $p\text{CO}_2$ exhibited substantial variability among the northern, central, and southern Bay of Bengal, which can be attributed to regional differences in water column mixing and biological productivity (Joshi & Warrior, 2022).

In the Northern BoB, the highest $p\text{CO}_2$ levels ($488.78 \mu\text{atm}$) were recorded during the post-monsoon season (Fig. 5). This observation is consistent with earlier reports by Sarma et al. (2012) and Shaik et al. (2023), which suggest that intense precipitation during the preceding monsoon reduces both salinity and temperature, thereby enhancing CO_2 solubility and resulting in elevated surface $p\text{CO}_2$. Since CO_2 solubility is a key factor regulating the regional carbon sink, this process plays a critical role in shaping the seasonal variability of surface $p\text{CO}_2$. Elevated values during February can also be linked to northeast monsoon winds, which promote upwelling and transport CO_2 -enriched subsurface waters to the surface, increasing $p\text{CO}_2$ concentrations (Joshi & Warrior, 2022). Conversely, the lowest $p\text{CO}_2$ levels ($248.75 \mu\text{atm}$) were observed in May (summer), largely due to elevated sea surface temperatures that reduce CO_2 solubility (Sarma et al., 2012; Sridevi & Sarma, 2021). In addition to hydrographic drivers, surface $p\text{CO}_2$ in this region is influenced by biological productivity, which is regulated by nutrient inputs from riverine fluxes (Kumar et al., 1996). The Ganges–Brahmaputra river system serves as a major nutrient source, stimulating primary production. Furthermore, the East India Coastal Current (EICC) flows southward from February to May, introducing high-salinity waters into the coastal region and promoting stratification, which is associated with reduced surface $p\text{CO}_2$ (Sarma et al., 2018).

In the central Bay of Bengal (BoB), $p\text{CO}_2$ concentrations ranged from $215.88 \mu\text{atm}$ in May, 2022 to $394.4 \mu\text{atm}$ in January, 2022, indicating generally lower values compared to the northern and southern BoB (Fig. 5). This reduction is strongly influenced by the seasonal reversal of the East India Coastal Current (EICC), which alternates between equatorward and poleward flow, transporting low- and high-salinity waters. The resulting surface stratification suppresses vertical mixing, thereby contributing to the overall lower $p\text{CO}_2$ trend across the region (Rao et al., 1994; Naqvi et al., 1994; Kumar et al., 1996; Shetye et al., 1993; Sanilkumar et al., 1997; Babu et al., 2003; Gangopadhyay et al., 2013; Sarma et al., 2013). Elevated $p\text{CO}_2$ values, however, were observed during the monsoon seasons, when northeast and southwest monsoon winds promote upwelling and bring CO_2 -rich subsurface waters to

the surface (Shetye et al., 1996). The lowest $p\text{CO}_2$ levels, recorded in April, are associated with the southward transport of low-salinity waters by the EICC, coupled with Ekman pumping (Hongyu et al., 2022). The presence of fresher water enhances CO_2 solubility, which, as suggested by Shaik (2023), contributes to reduced surface $p\text{CO}_2$. Moreover, during April (summer), elevated sea surface temperatures further influence salinity. As highlighted by Vajravelu et al. (2018), higher temperatures increase evaporation and salinity, disrupting CO_2 solubility and reinforcing the low $p\text{CO}_2$ conditions observed during this period.

In the southern BoB, $p\text{CO}_2$ values ranged from 217.51 μatm (April–May, 2022) to 372.89 μatm (January, post-monsoon 2023), following a seasonal pattern similar to that observed in the central BoB (Fig. 5). Seasonal hydrographic conditions, coupled with variable riverine inputs influenced by the incursion of the Southwest Monsoon Current (SMC), play a vital role in regulating productivity in this region. Enhanced nutrient supply promotes primary production, which subsequently increases biogenic carbon flux to the southern BoB (Ittekkot, 1991). Consistent with observations by Naveen et al. (2023), $p\text{CO}_2$ minima generally occurred during summer, while maxima were recorded during the monsoon season. Elevated $p\text{CO}_2$ during the monsoon can be attributed to the influx of organic and inorganic matter from terrestrial sources delivered through riverine discharge. These inputs stimulate heterotrophic bacterial respiration, releasing CO_2 into subsurface waters (Kempe et al., 1991; Cai & Wang, 1998). The CO_2 -enriched subsurface waters are then transported to the surface via physical and biological processes, including nutrient upwelling driven by strong monsoonal winds (Sarma et al., 1996). Conversely, minimum $p\text{CO}_2$ levels during summer are associated with elevated sea surface temperatures that enhance stratification, thereby restricting the upward transport of CO_2 -rich subsurface waters (Kumar et al., 2007; Sarma et al., 2000; Takahashi, 2002).

5. Conclusion

The present study highlights the spatial and temporal variability of surface $p\text{CO}_2$ across Northern, Central and Southern BoB during 2022 – 2024. Distinct seasonal patterns observed, with consistently higher $p\text{CO}_2$ observed with post-monsoon and pre-monsoon season and diminished levels during summer across entire BoB. The higher $p\text{CO}_2$ levels in Northern BoB attributed to the riverine influx from Ganges-Brahmaputra region, reduction in salinity diminishes the CO_2 solubility and monsoon-driven upwelling. In contrast, Central BoB exhibited a comparatively low $p\text{CO}_2$ levels primarily regulated by EICC current, which

takes control of salinity and stratification. The southern BoB too exhibits a similar trend with Central BoB but the main responsible factors are river inputs and upwelling process. In regard with the outcome of the present study suggests that BoB function as dynamic and heterogenous system, where regional hydrography and seasonal monsoon variations governs that carbon dynamics. Long term monitoring of carbon dynamics, satellite observation found vital to understand the BoB's role in global carbon cycle under increasing climate change.

6. References

- Akhand, A., Chanda, A., Watanabe, K., Das, S., Tokoro, T., Hazra, S., Kuwae, T., 2022. Drivers of inorganic carbon dynamics and air–water CO₂ fluxes in two large tropical estuaries: Insights from coupled radon (²²²Rn) and pCO₂ surveys. *Limnology and Oceanography* 67, S118–S132
- Anderson, T. R. (2005). Plankton functional type modelling: running before we can walk?. *Journal of plankton research*, 27(11), 1073-1081.
- Anderson, T. R. (2010). Progress in marine ecosystem modelling and the “unreasonable effectiveness of mathematics”. *Journal of Marine Systems*, 81(1-2), 4-11.
- Babu MT, Sarma YVB, Murty VSN, Vethamony P (2003) On the circulation in the Bay of Bengal during Northern spring inter-monsoon (March-April 1987). *Deep-Sea Res Part II: Topical Stud Oceanography* 50:5. [https://doi.org/10.1016/S0967-0645\(02\)00609-4](https://doi.org/10.1016/S0967-0645(02)00609-4).
- Blomquist, B. W., Brumer, S. E., Fairall, C. W., Huebert, B. J., Zappa, C. J., Brooks, I. M., ... & Pascal, R. W. (2017). Wind speed and sea state dependencies of air-sea gas transfer: Results from the high wind speed gas exchange study (HiWinGS). *Journal of Geophysical Research: Oceans*, 122(10), 8034-8062.
- Borges, A. V. (2005). Do we have enough pieces of the jigsaw to integrate CO₂ fluxes in the coastal ocean?. *Estuaries*, 28(1), 3-27.
- Cai, B., Zhu, J., Ban, F., & Tan, M. (2011). Intra-annual variation of the calcite deposition rate of drip water in Shihua Cave, Beijing, China and its implications for palaeoclimatic reconstructions. *Boreas*, 40(3), 525-535.
- Cai, W. J., & Wang, Y. (1998). The chemistry, fluxes, and sources of carbon dioxide in the estuarine waters of the Satilla and Altamaha Rivers, Georgia. *Limnology and Oceanography*, 43(4), 657-668.
- Dai, M., Cao, Z., Guo, X., Zhai, W., Liu, Z., Yin, Z., ... & Du, C. (2013). Why are some marginal seas sources of atmospheric CO₂?. *Geophysical Research Letters*, 40(10), 2154-2158.
- Doney, S. C., Balch, W. M., Fabry, V. J., & Feely, R. A. (2009). Ocean acidification: a critical emerging problem for the ocean sciences. *Oceanography*, 22(4), 16-25.
- Doney, S. C., Mahowald, N., Lima, I., Feely, R. A., Mackenzie, F. T., Lamarque, J. F., & Rasch, P. J. (2007). Impact of anthropogenic atmospheric nitrogen and sulfur deposition on

ocean acidification and the inorganic carbon system. *Proceedings of the National Academy of Sciences*, 104(37), 14580-14585.

Gangopadhyay, A., Bharat Raj, G. N., Chaudhuri, A. H., Babu, M. T., & Sengupta, D. (2013). On the nature of meandering of the springtime western boundary current in the Bay of Bengal. *Geophysical Research Letters*, 40(10), 2188-2193.

Hall, E. R., Wickes, L., Burnett, L. E., Scott, G. I., Hernandez, D., Yates, K. K., ... & Styron, J. (2020). Acidification in the US Southeast: Causes, potential consequences and the role of the southeast ocean and coastal acidification network. *Frontiers in Marine Science*, 7, 548.

Ho, D. T., Law, C. S., Smith, M. J., Schlosser, P., Harvey, M., & Hill, P. (2006). Measurements of air-sea gas exchange at high wind speeds in the Southern Ocean: Implications for global parameterizations. *Geophysical Research Letters*, 33(16).

Hongyu, X. I. N., Qiang, X. I. E., & Weiqiang, W. A. N. G. (2022). Seasonal variation of East India Coastal Current and its transports of heat and salt. *Journal of Tropical Oceanography*, 41(2), 38-51.

Ittekkot, V., Nair, R. R., Honjo, S., Ramaswamy, V., Bartsch, M., Manganini, S., & Desai, B. N. (1991). Enhanced particle fluxes in Bay of Bengal induced by injection of fresh water. *Nature*, 351(6325), 385-387.

Joshi, A. P., & Warrior, H. V. (2022). Comprehending the role of different mechanisms and drivers affecting the sea-surface pCO₂ and the air-sea CO₂ fluxes in the Bay of Bengal: A modeling study. *Marine Chemistry*, 243, 104120.

Joshi, A. P., Chowdhury, R. R., Warrior, H. V., & Kumar, V. (2021). Influence of the freshwater plume dynamics and the barrier layer thickness on the CO₂ source and sink characteristics of the Bay of Bengal. *Marine Chemistry*, 236, 104030.

KEMPE, S., Pettine, M., & CAUWET, G. (1991). Biogeochemistry of European rivers. In *Biogeochemistry of major world rivers. SCOPE 42* (pp. 169-211).

Kumar, M. D., Naqvi, S. W. A., George, M. D., & Jayakumar, D. A. (1996). A sink for atmospheric carbon dioxide in the northeast Indian Ocean. *Journal of Geophysical Research: Oceans*, 101(C8), 18121-18125.

Kumar, S. P., Nuncio, M., Ramaiah, N., Sardesai, S., Narvekar, J., Fernandes, V., & Paul, J. T. (2007). Eddy-mediated biological productivity in the Bay of Bengal during fall and spring intermonsoons. *Deep Sea Research Part I: Oceanographic Research Papers*, 54(9), 1619-1640.

Laruelle, G. G., Cai, W. J., Hu, X., Gruber, N., Mackenzie, F. T., & Regnier, P. (2018). Continental shelves as a variable but increasing global sink for atmospheric carbon dioxide. *Nature communications*, 9(1), 454.

Le Qu' er' e, C., Andrew, R.M., Friedlingstein, P., et al., 2018. Global Carbon Budget 2018. *Earth Syst. Sci. Data* 10, 2141–2194. <https://doi.org/10.5194/essd-10-2141-2018>. Sabine, C.L., Feely, R.A., Gruber, N., et al., 2004. The Oceanic Sink for Anthropogenic CO₂. *Science* 305, 367–371. <https://doi.org/10.1126/science.1097403>.

- Li, Q., Guo, X., Zhai, W., Xu, Y., & Dai, M. (2020). Partial pressure of CO₂ and air-sea CO₂ fluxes in the South China Sea: Synthesis of an 18-year dataset. *Progress in Oceanography*, 182, 102272.
- Liu, J., Bellerby, R.G.J., Li, X., Yang, A., Yang. 2022. Seasonal variability of the carbonate system and air-sea CO₂ flux in the outer Changjiang Estuary, East China Sea. *Frontiers in Marine Science* 8, 765564
- Mackenzie, T., Ver, L. M., & Lerman, A. (1995). Modeling the behaviour of biogeochemical cycles in the land-coastal margin system during global change. In *International Symposium on Earth Environment* (pp. 149-200).
- Muraleedharan, P. M., & Prasannakumar, S. (1996). Arabian Sea upwelling-A comparison between coastal and open ocean regions.
- Naqvi, S. W. A., Jayakumar, D. A., Nair, M., Kumar, M. D., & George, M. D. (1994). Nitrous oxide in the western Bay of Bengal. *Marine chemistry*, 47(3-4), 269-278
- Naveen, M., Priyanka, K., Shanthi, R., Utthamapandiyar, U., Saravanakumar, A., Roy, R., & Nagamani, P. V. (2023). Spatial and temporal variability of the sources and sinks of carbonate system in the southwest bay of Bengal from 2014 to 2020. *Quaternary Science Advances*, 11, 100080.
- Nightingale, P. D., Malin, G., Law, C. S., Watson, A. J., Liss, P. S., Liddicoat, M. I., ... & Upstill-Goddard, R. C. (2000). In situ evaluation of air-sea gas exchange parameterizations using novel conservative and volatile tracers. *Global Biogeochemical Cycles*, 14(1), 373-387.
- Rao, C. K., Naqvi, S. W. A., Kumar, M. D., Varaprasad, S. J. D., Jayakumar, D. A., George, M. D., & Singbal, S. Y. S. (1994). Hydrochemistry of the Bay of Bengal: possible reasons for a different water-column cycling of carbon and nitrogen from the Arabian Sea. *Marine Chemistry*, 47(3-4), 279-290.
- Sabine, C.L., Feely, R.A., Gruber, N., et al., 2004. The Oceanic Sink for Anthropogenic CO₂. *Science* 305, 367–371. <https://doi.org/10.1126/science.1097403>. Sarmiento, J.L., Gruber, N., 2006. *Ocean Biogeochemical Dynamics*. Princeton Univ. Press
- Sanilkumar, K. V., Kuruvilla, T. V., Jogendranath, D., & Rao, R. R. (1997). Observations of the Western Boundary Current of the Bay of Bengal from a hydrographic survey during March 1993. *Deep Sea Research Part I: Oceanographic Research Papers*, 44(1), 135-145.
- Sarma, V. V. S. S., Krishna, M. S., Paul, Y. S., & Murty, V. S. N. (2015). Observed changes in ocean acidity and carbon dioxide exchange in the coastal Bay of Bengal—a link to air pollution. *Tellus B: Chemical and Physical Meteorology*, 67(1), 24638.
- Sarma, V. V. S. S., Kumar, G. S., Yadav, K., Dalabehera, H. B., Rao, D. N., Behera, S., & Loganathan, J. (2019). Impact of eddies on dissolved inorganic carbon components in the Bay of Bengal. *Deep Sea Research Part I: Oceanographic Research Papers*, 147, 111-120.
- Sarma, V. V. S. S., Kumar, M. D., George, M. D., & Rajendran, A. (1996). Seasonal variations in inorganic carbon components in the central and eastern Arabian Sea. *Current Science*, 852-856.

Sarma, V. V. S. S., Kumari, V. R., Srinivas, T. N. R., Krishna, M. S., Ganapathi, P., & Murty, V. S. N. (2018). East India Coastal Current controls the dissolved inorganic carbon in the coastal Bay of Bengal. *Marine Chemistry*, 205, 37-47.

Sarma, V. V. S. S., Sridevi, B., Maneesha, K., Sridevi, T., Naidu, S. A., Prasad, V. R., Venkataramana, V., Acharya, T., Bharati, M. D., Subbaiah, C. V., Kiran, B. S., Reddy, N p. C., Sarma, V. V., Sadhuram, Y., & Murty, T. V. R. (2013). Impact of atmospheric and physical forcings on biogeochemical cycling of dissolved oxygen and nutrients in the coastal Bay of Bengal. *Journal of oceanography*, 69, 229-243.

Sarma, V. V. S. S., Viswanadham, R., Rao, G. D., Prasad, V. R., Kumar, B. S. K., Naidu, S. A., kumar, N.A., Rao, D. B., Sridevi, T., krishna, M. S., Reddy, N. P. C., Sadhuram, Y., & Murty, T. V. R. (2012). Carbon dioxide emissions from Indian monsoonal estuaries. *Geophysical Research Letters*, 39(3).

Sarma, V.V.S.S., Kumar, M.D., Gauns, M., Madhupratap, M., 2000. Seasonal controls on surface pCO₂ 471–479.

Schlesinger, W. H. (1997). Biogeochemistry: an analysis of global change.

Shaik, I., Krishna, K. V., Begum, S. K., Suhail, K. M., Nagamani, P. V., Shanmugam, P., & Srivathsav, K. (2023). Advancements in Carbon Dioxide Modeling: An Algorithm Incorporating *In-situ* and Satellite Data for Improved Understanding of pCO Dynamics in the Bay of Bengal. *IEEE Access*, 11, 144877-144886.

Shanthi, R., Poornima, D., Naveen, M., Thangaradjou, T., Choudhury, S. B., Rao, K. H., & Dadhwal, V. K. (2016). Air-sea CO₂ flux pattern along the southern Bay of Bengal waters. *Dynamics of Atmospheres and Oceans*, 76, 14-28.

Shen, C., Testa, J.M., Li, M., Chen, B., Cai, W.J., 2023. Interannual variability of air-water CO₂ flux in a large eutrophic estuary. *Water Research* 244, 120523.

Shetye, S. R., Gouveia, A. D., Shankar, D., Shenoi, S. S. C., Vinayachandran, P. N., Sundar, D., Michael, G. S., & Nampoothiri, G. (1996). Hydrography and circulation in the western Bay of Bengal during the northeast monsoon. *Journal of Geophysical Research: Oceans*, 101(C6), 14011-14025.

Shetye, S. R., Gouveia, A. D., Shenoi, S. S. C., Sundar, D., Michael, G. S., & Nampoothiri, G. (1993). The western boundary current of the seasonal subtropical gyre in the Bay of Bengal. *Journal of Geophysical Research: Oceans*, 98(C1), 945-954.

Sridevi, B., & Sarma, V. V. S. S. (2021). Role of river discharge and warming on ocean acidification and pCO₂ levels in the Bay of Bengal. *Tellus B: Chemical and Physical Meteorology*, 73(1), 1-20.

Takahashi, T., Sutherland, S. C., Sweeney, C., Poisson, A., Metzl, N., Tilbrook, B., & Nojiri, Y. (2002). Global sea–air CO₂ flux based on climatological surface ocean pCO₂, and seasonal biological and temperature effects. *Deep Sea Research Part II: Topical Studies in Oceanography*, 49(9-10), 1601-1622.

Takahashi, T., Sutherland, S.C., Wanninkhof, R., et al., 2009. Climatological mean and decadal change in surface ocean pCO₂, and net sea–air CO₂ flux over the global oceans. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 56, 554–577. <https://doi.org/10.1016/j.dsr2.2008.12.009>.

Unesco. (1969). *Discharge of selected rivers of the world*. Unesco.

Vajravelu, M., Martin, Y., Ayyappan, S., & Mayakrishnan, M. (2018). Seasonal influence of physico-chemical parameters on phytoplankton diversity, community structure and abundance at Parangipettai coastal waters, Bay of Bengal, South East Coast of India. *Oceanologia*, 60(2), 114-127.

Wanninkhof, R. (1992). Relationship between wind speed and gas exchange over the ocean. *Journal of Geophysical Research: Oceans*, 97(C5), 7373-7382.

Wanninkhof, R., Asher, W. E., Ho, D. T., Sweeney, C., & McGillis, W. R. (2009). Advances in quantifying air-sea gas exchange and environmental forcing. *Annual review of marine science*, 1(1), 213-244.

Zhai, W. D., Dai, M. H., Chen, B. S., Guo, X. H., Li, Q., Shang, S. L., ... & Wang, D. X. (2013). Seasonal variations of sea-air CO₂ fluxes in the largest tropical marginal sea (South China Sea) based on multiple-year underway measurements. *Biogeosciences*, 10(11), 7775-7791.

Author Contributions: The first draft of the manuscript was written by Lowin Anand and supervision by Ayyapan Saravankumar. All authors read and approved the final manuscript.

Acknowledgement:

We acknowledge the authorities of CAS in Marine Biology, Faculty of Marine Sciences, Annamalai University, Parangipettai, Tamil Nadu, India, and NRSC-pCO₂ for all the support to carry out the work.

Declaration of Interest statement:

Funding: This work was supported by NRSC-pCO₂

Competing Interest: The authors have no relevant financial or non-financial interests to disclose

Ethics approval: Not applicable

Constant to Publish: Not applicable

Conflict of Interest Statement: None declared

Data availability statement: Data available on request